

# EMPIRICALLY-CONSTRAINED CLIMATE SENSITIVITY AND THE SOCIAL COST OF CARBON

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DRAFT FOR COMMENTS: April 5, 2016

**Abstract:** Integrated Assessment Models (IAMs) require parameterization of both economic and climatic processes. The latter include *Ocean Heat Uptake* (OHU) efficiency, which represents the rate of heat exchange between the atmosphere and the deep ocean, and *Equilibrium Climate Sensitivity* (ECS), or the surface temperature response to doubling of CO<sub>2</sub> levels after adjustment of the deep ocean. Due to a lack of adequate data, OHU and ECS parameter distributions in IAMs have been based on simulations from climate models. In recent years, new and sufficiently long observational data sets have emerged to support a growing body of empirical ECS estimates, but the results have not been applied in IAMs. We incorporate a recent observational estimate of the ECS distribution conditioned on observed OHU efficiency into two widely-used IAMs. The resulting Social Cost of Carbon (SCC) estimates are much smaller than those from models based on simulated parameters. In the DICE model the average SCC falls by 30-50% depending on the discount rate, while in the FUND model the average SCC falls by over 80%. The span of estimates across discount rates also shrinks considerably, implying less sensitivity to this parameter choice.

Keywords: Social Cost of Carbon, Climate Sensitivity, Carbon Taxes, Integrated Assessment Models

[1]

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## 1 INTRODUCTION

Integrated Assessment Models (IAMs) are central to the analysis of climate policy, especially for estimating the social cost of carbon (SCC), i.e. the discounted present value of the marginal damages of a tonne of carbon dioxide emissions. IAMs operate at a high level of abstraction and require extensive parameterization for both climatic and economic processes. Two key climate parameters are the equilibrium climate sensitivity (ECS), representing the long term temperature change from doubling atmospheric CO<sub>2</sub>, and the ocean heat uptake (OHU) efficiency, representing the rate at which the ocean sequesters atmospheric heat. Among the economic parameters, the most influential are the discount rate and the coefficients of the Marginal Damages function (Marten 2011).

Marginal damages are, in turn, strongly influenced by the way ECS is represented. The customary way is through use of a probability density function (PDF) parameterized to fit a range of estimates from climate modeling simulations, which then gives rise to a PDF of marginal damages. The use of an ECS distribution rather than a point estimate can strongly influence the average SCC if it has a large upper tail, which pulls up both the median and mean SCC values. A commonly-used ECS PDF is taken from a graph in Roe and Baker (2007, herein RB07) which does have a long upper tail. RB07 was an exploration of why uncertainties over ECS have not been

reduced despite decades of effort, with the explanation centering on the amplified effect of uncertainties in the value of the climate feedback parameter  $f$  on final temperatures, due to its position in the denominator of the equation for ECS. To illustrate the point they fitted a curve to a small selection of ECS estimates published between 2003 and 2007, yielding an ECS curve that had a long upper tail even though there was no unbounded source of uncertainty in the underlying model.

IAMs have relied heavily on the RB07 graph to characterize ECS, most notably for the purpose of the highly-influential reports of the US Interagency Working Group (IWG 2010, 2013) which determines the SCC for official US regulatory purposes. But this usage is inappropriate for two reasons. First, as Roe himself later pointed out (Roe and Bauman 2013), the distribution in RB07 was not applicable in the context of IAM simulations because the wideness of the tails is a function of the time span to equilibrium, and the time span relevant to IAM simulations is not consistent with a fat tailed-ECS distribution. In the real world, CO<sub>2</sub> doubling is not instantaneous, the transition to a new equilibrium state is exceedingly slow, and the oceans absorb huge amounts of heat along the way. In simplified climate models, time-to-equilibrium goes up with the square of ECS, so an upward adjustment of the ECS parameter without taking into account the slower time path of warming will yield distorted present value damage estimates (Roe and Bauman 2013). In particular, the higher the ECS, the slower the adjustment process, making the fat upper tail of realized warming physically impossible for even a thousand years into the future (Roe and Bauman 2013, p. 653). Thus, under even minuscule discount rates, the fatness of the tail in RB07 is not relevant for SCC calculations and its usage in probabilistic IAM simulations generates a misleading distribution of SCC values. An ECS distribution applicable to the real world must therefore be

constrained by a realistic OHU efficiency estimate. Roe and Bauman (2013) illustrate this by recalculating the Weitzman (2009) insurance model under the fat-tails ECS assumption, showing that applying the relevant physical constraints cuts the value of insuring against catastrophe to below 1 percent of GDP.

Second, RB07 predated a large literature on empirical ECS estimation. As was common at the time, they fitted a distribution to a small number of simulated ECS distributions derived from climate models. It is only relatively recently that sufficiently long and detailed observational data sets have been produced to allow direct estimation of ECS using empirical energy balance models. A large number of studies have appeared since 2010 estimating ECS on long term climatic data (Otto et al. 2013, Ring et al. 2012, Aldrin et al. 2012, Lewis 2013, Lewis & Curry 2015, Schwartz 2012, Skeie et al 2014, Lewis 2016, etc.). This literature has consistently yielded ECS values near or even below the low end of the range taken from climate model studies. General circulation models (GCMs) historically yielded sensitivities in the range of 2.0 – 4.5 °C, and (based largely on GCMs) RB07 yields a central 90 percent range of 1.72 – 7.14 °C with a median of 3.0 °C and a mean of 3.5 °C (see comparison table in IWG 2010, p. 13). But the median of recent empirical estimates has generally been between 1.5 and 2.0 °C, with 95% uncertainty bounds below the RB07 average. While this inconsistency has attracted growing attention in the climatology literature (Kummer and Kessler 2014, Marvel et al. 2015) IAM practitioners have ignored it and use only the older GCM-based parameterizations in their models.

Both these concerns point to a potential for existing SCC calculations to be biased too high. We investigate this by replicating standard SCC estimates using two leading IAMs (FUND and DICE), then re-doing the calculations using an observational ECS distribution from a recent study (Lewis

and Curry 2015, herein LC15) that controls for the effect of ocean heat sequestration, thereby yielding an empirically-constrained climate sensitivity distribution. We find that the resulting SCC values drop dramatically compared to those reported in the IWG (2010, 2013). The model-based RB07 ECS distribution at a 3 percent discount rate yields a mean SCC for the year 2020 of \$37.73, in line with the IWG estimates that currently guide US policymaking. Substituting an empirical ECS distribution from LC15 yields a mean 2020 SCC of \$19.52, a drop of 48%. The same exercise for the FUND model yields a mean SCC estimate of \$19.33 based on RB07 and \$3.33 based on the LC15 parameters—an 83% decline. Furthermore the probability of a negative SCC (implying CO<sub>2</sub> emissions are a positive externality) jumps dramatically using an empirical ECS distribution. Using the FUND model, under the RB07 parameterization at a 3% discount rate there is only about a ten percent chance of a negative SCC through 2050, but using the LC15 distribution, the probability of a negative SCC jumps to about 40%. Remarkably, replacing simulated climate sensitivity values with an empirical distribution calls into question whether CO<sub>2</sub> is even a negative externality. The lower SCC values also cluster more closely together across different discount rates, diminishing the importance of this parameter.

The paper proceeds as follows. Section 2 explains the roles of ECS and OHU parameterization in climate submodels, and reviews the empirical literature over the past half-decade. Section 3 presents SCC calculations using DICE and FUND, and Section 4 presents conclusions.

## **2 IAM PARAMETERIZATION**

ECS is defined as the average increase in temperatures around the world as a result of CO<sub>2</sub> doubling, after the deep ocean has adjusted to the increased forcing. While data on historical

temperatures and CO<sub>2</sub> concentrations are available, it is not straightforward to estimate ECS. Rising levels of CO<sub>2</sub> and other greenhouse gases must be translated into units of “radiative forcing” which maps their effect on radiation into a common measure by which the effects of all types of drivers of climate change on temperature change may be compared. The warming effect of CO<sub>2</sub> and other greenhouse gases is partially offset by the potential cooling effect of aerosols which are sometimes released by the same processes responsible for CO<sub>2</sub>. However, compared to the warming effect of CO<sub>2</sub>, the direct and indirect aerosol cooling effects, and the related negative aerosol forcing, are much more uncertain and difficult to quantify, in part because of their interactions with cloud formation (IPCC 2013 ch. 8, Kiehl 2007, Schwartz et al 2007). Hence there is a range of possible forcing values consistent with the historical CO<sub>2</sub> record, based on how strong the offsetting aerosol forcing is taken to be.

If cooling by aerosol forcing is strongly negative, it will offset much of the positive, warming, forcing from greenhouse gases; if it is weak it will offset little of it. The net forcing, in turn, affects temperature according to the magnitude of ECS. Since the historical temperature record is fixed, there must be an offsetting relationship between ECS and estimated forcing: for a given temperature change, greater net historical forcing implies lower ECS and vice-versa. This inverse relationship is reflected across the suite of climate models. For instance, models that translate historical greenhouse gas and aerosol levels into a relatively strong positive forcing must have lower ECS, etc. (Kiehl 2007).

The treatment within a model of OHU efficiency also affects the ECS that corresponds to observed warming. The higher OHU efficiency is, and thus the larger the amount of heat sequestered in the oceans over the past century, the more the historical climate record understates

the total amount of warming that will ultimately occur. However, considering the massive heat capacity of the ocean this also implies that future warming will be likewise sequestered relatively efficiently, reducing the upper tail of ECS estimates (Roe and Bauman 2013).

Consequently, estimates of ECS need to make use of information on aerosol forcing and OHU efficiency as well as CO<sub>2</sub> and temperature records. The problem with calibrating an IAM using an ECS curve fitted to data generated by unconstrained climate models is that the range of values is not necessarily consistent with the world as represented in the IAM. The spread of model ECS measures taken in isolation may be unrealistic and would certainly be misleading if applied in an IAM without being conditioned on realistic OHU efficiency estimates. In other words the range of ECS values in ensembles of GCM runs is not empirically-constrained. One reason for this is that many processes critical to ECS and OHU values in a GCM are sub-grid scale and must be represented by crude parameterised approximations that are not themselves constrained by observation. Zhao et al (2016), for instance, shows that a wide range of ECS values can emerge in GCMs by altering a single cloud microphysics parameter, but since the models all exhibit roughly equal fidelity to historical observations there is no way of telling which value is most appropriate. Consequently when IWG (2010, p. 13) reports that there is only a 1.3 percent probability that ECS is less than 1.5 °C, this should be understood as a statement about the ensemble of GCMs they were considering, but it is not necessarily a statement about the real world.

An alternative approach that has a somewhat better chance of yielding statements applicable to the real world involves estimating an empirical energy balance model using aerosol forcing and historical ocean heat content estimates to condition the ECS distribution. This is the approach taken in LC15. They used the 1750-2011 forcing and OHU estimates from the then-most recent IPCC

report (IPCC 2014), yielding a median ECS of 1.64 °C and a 5—95 % uncertainty range of 1.05 – 4.05 °C. This is in line with empirical estimates from Otto et al. (2013), Ring et al. (2012), Aldrin et al. (2012) and Lewis (2013), but is in clear contrast to the customary IAM parameterization using RB07, since the central value in LC15 falls below the 5% lower bound of the ECS distribution used in IWG (2010, 2013). Not surprisingly, this implies that empirically unconstrained SCC estimates are skewed high.

One proposal for resolving the discrepancy between model-based and empirical ECS estimates is to use so-called “efficacies” to adjust forcing values for different feedback responses (Kummer and Dessler, 2014, Marvel et al. 2015). The underlying argument is that two different types of climate forcing agents, each with equivalent initial forcing, may still generate different eventual temperature impacts, particularly if they are not well-mixed and the spatial variation induces differing local feedbacks. Hansen (2005) introduced the term “efficacy” to capture this concept, with the unit of measurement a ratio of overall temperature response to that of CO<sub>2</sub>. If a species of, say, greenhouse gas has an efficacy of 1.5 this means that if its atmospheric concentration were to increase by an amount corresponding to the same additional effective radiative forcing (ERF) as a doubling of CO<sub>2</sub> levels, the eventual temperature response would be 1.5 times larger, due to the different feedback processes involved. Hansen (2005) found in GCM simulations that most efficacies were close to unity so the spatial and other variations did not matter much at the global level. Marvel et al. (2015) analyse simulated data from a different model (GISS-E2-R) and found that aggregate efficacy of forcings operating over the historical period was below one, due *inter alia* to aerosol forcing (which is negative) having a high efficacy. They argued that this meant empirical estimates of ECS were biased down. However, Marvel et al. mistakenly left out land-use forcing, and

their revised uncertainty bounds are so wide that most single-forcing efficacies for ERF encompass unity (Marvel et al. 2016 Table S1), though the aggregate efficacy remained below unity. More generally, Marvel et al. (2015) is a study of the behavior of a particular GCM, and it has not been established whether forcing efficacy differences among climate models translates into empirically-relevant real-world effects. Otto et al. (2013), for instance, used forcing estimates that implicitly incorporated efficacy variations and they obtained an ECS distribution nearly identical to that in LC15.

### **3 SCC CALCULATIONS USING EMPIRICAL PARAMETERS**

We first replicated the SCC estimates that would have been used in IWG (2013) from both the DICE and FUND models based on the RB07 ECS distribution. As we did not include the PAGE model in our work (due to the unavailability of the code) we cannot directly compare our results with the IWG tables since they are averaged over all three models. IWG (2013) Table A5 lists separate results for FUND and DICE for 2020 and we were able to check our results against those. Since the calculations are probabilistic it is not guaranteed that we will reproduce the exact SCC estimates as shown in IWG (2013), but our replication is quite close. Table 1 shows the DICE and FUND SCC estimates for 2020 compared with our replications (“Repl”) for three discount rates. Apart from a slight under-estimation of the FUND results under the lower discount rates the match is extremely good.

#### **3.1 DICE MODEL**

Table 2 shows the mean SCC estimates for four discount rates, applying the RB07 and LC15 ECS distribution to the DICE model. The final row shows the percentage change for the 2020 estimates

(all years exhibit about the same percentage changes). Under the widely-used RB07 distribution, the SCC ranges from \$3.88 to \$89.26 depending on the discount rate and the future year. Under the LC15 parameter distributions the SCC ranges from \$2.39 to \$46.00. The largest proportional drop—nearly 50 percent—is observed in the low discount rate case. The high discount rate case yields a drop of just under 40 percent.

These reductions are primarily due to the LC15 distribution containing a smaller upper tail and therefore greater probability mass at lower temperatures. Table 3 shows the average standard deviations of the two sets of estimates. The largest reduction, about 24 percent, again occurs at the lowest discount rate, compared to only seven percent at the highest discount rate. The LC15 distribution provides uniformly more certainty for the SCC for all years and all discount rates. These results are in line with previous research performing similar computations by applying the Otto et al (2013) ECS distribution in the DICE model (Dayaratna and Kreutzer 2013).

### 3.2 FUND MODEL

Tables 4 and 5 present the same results as Tables 2 and 3, but for the FUND model. A number of differences are notable. The mean SCC estimates are lower under both parameterizations, and under the empirical LC15 coefficients are, on average, negative at 5 percent or higher discount rates out past 2030. This implies that carbon dioxide emissions are a positive externality, so that the optimal policy would require subsidizing emissions. Also, in contrast to the DICE model, use of the LC15 coefficients increases the average standard deviation, indicating higher uncertainty. The increased uncertainty includes a much larger lower tail, implying a larger probability of a negative SCC. Table 6 shows that, under the RB07 parameterization, at a three percent discount rate the

probability of carbon dioxide emissions being a positive externality is between nine and 12 percent out to 2050. But using the LC15 parameters this probability jumps to between 37 and 45 percent.

These results are in line with previous simulations using other ECS distributions that have smaller upper tails than RB07, namely Otto et al (2013) and Lewis (2013); see Dayaratna and Kreutzer (2014).

#### **4 DISCUSSION AND CONCLUSION**

Model-based ECS distributions are misleading for use in SCC calculations because they are skewed high relative to abundant empirical evidence and because of their lack of constraint to OHU efficiency rates relevant to IAM timelines, which yields a physically-unrealistic upper tail. The model-observational mismatch in ECS estimation is not attributable simply to a specific empirical methodology, as very similar results have been found by Otto et al. (2013), Ring et al. (2012), Aldrin et al. (2012) and others using a variety of methods. Nor is it an artifact of selecting a specific estimation period, as LC15 showed their results were robust to numerous variations on the choice of base and final periods (LC15, Table 4).

We incorporated the Lewis and Curry (2015) ECS distribution, which is conditioned on updated forcings and OHU data, into the DICE and FUND models. This reduces the estimated Social Cost of Carbon in both, regardless of discount rates. Using a 3 percent discount rate and the RB07 ECS distribution, DICE yields an average SCC ranging from about \$30 to \$60 between now and 2050, but this falls in half to the \$15 to \$30 range using the LC15 ECS estimate. The corresponding average SCC in FUND falls from the \$17 to \$27 range to the \$3 to \$5 range. Moreover FUND, which takes more explicit account of potential regional benefits from CO<sub>2</sub> fertilization and increased agricultural

productivity, yields a substantial (about 40 percent or more) probability of a negative SCC through the first half of the 21<sup>st</sup> century.

A further way in which use of empirically-constrained parameters reduces uncertainty is the shrinking of the SCC range across discount rates. In the DICE model under the RB07 parameterization, the mean SCC estimates span about \$50 as of 2010 depending on choice of discount rate, with the span rising to about \$80 as of 2050. This span shrinks to the \$20 to \$40 range under the LC15 parameterization. Using the FUND model, the uncertainty range associated with the choice of discount rate is from about \$30 to \$40 under the RB07 parameterization, falling to \$4 to \$8 range under the LC15 parameterization. Thus, use of well-constrained empirical parameters makes a substantial contribution also to reducing uncertainty associated with the choice of discount rate.

## 5 ACKNOWLEDGMENTS

We thank Nicholas Lewis for comments on an earlier draft. Funding support from the Heritage Foundation, the Cato Institute and the Frontier Centre for Public Policy is gratefully acknowledged.

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## 7 TABLES

	2.5%		3.0%		5.0%	
	IWG	Repl	IWG	Repl	IWG	Repl
DICE	\$57	\$57	\$38	\$38	\$12	\$12
FUND	\$36	\$33	\$21	\$19	\$3	\$3

**Table 1:** Replication of IWG (2013) SCC estimates for DICE and FUND models for 2020, under three discount rate assumptions. Replications done herein denoted “Repl”.

Mean Social Cost of Carbon - DICE Model								
Discount rates	Using Simulated ECS				Using Empirical ECS			
	2.50%	3.00%	5.00%	7.00%	2.50%	3.00%	5.00%	7.00%
2010	\$46.87	\$29.96	\$8.62	\$3.88	\$23.72	\$15.50	\$4.91	\$2.39
2020	\$57.34	\$37.73	\$11.85	\$5.66	\$29.07	\$19.52	\$6.70	\$3.43
2030	\$67.13	\$45.18	\$15.04	\$7.43	\$34.20	\$23.45	\$8.48	\$4.47
2040	\$77.82	\$53.45	\$18.74	\$9.53	\$39.86	\$27.84	\$10.56	\$5.71
2050	\$89.26	\$62.38	\$22.90	\$11.94	\$46.00	\$32.65	\$12.90	\$7.13
% Chg at 2020					-49.3%	-48.3%	-43.5%	-39.3%

**Table 2:** Mean Social Cost of Carbon estimates by year under four discount rates from the DICE Model, for both the simulated (RB07) and empirical (LC15) ECS distributions. Last row shows the percent change as of 2020.

Average Standard Deviation - DICE Model								
Discount rates	Using Simulated ECS				Using Empirical ECS			
	2.50%	3.00%	5.00%	7.00%	2.50%	3.00%	5.00%	7.00%
2010	\$25.73	\$14.88	\$3.28	\$1.15	\$19.34	\$11.36	\$2.72	\$1.09
2020	\$31.29	\$18.74	\$4.59	\$1.75	\$23.74	\$14.32	\$3.78	\$1.62
2030	\$36.64	\$22.62	\$6.00	\$2.41	\$27.97	\$17.17	\$4.83	\$2.17
2040	\$42.22	\$27.00	\$7.79	\$3.25	\$32.61	\$20.26	\$6.06	\$2.86
2050	\$47.71	\$31.00	\$9.82	\$4.25	\$37.51	\$23.64	\$7.50	\$3.64
% Chg at 2020					-24.2%	-23.6%	-17.6%	-7.4%

**Table 3:** Average standard deviation of SCC estimates by year under four discount rates from the DICE Model, for both the simulated (RB07) and empirical (LC15) ECS distributions. Last row shows the percent change as of 2020.

Mean Social Cost of Carbon – FUND Model								
Discount rates	Using Simulated ECS				Using Empirical ECS			
	2.50%	3.00%	5.00%	7.00%	2.50%	3.00%	5.00%	7.00%
2010	\$29.69	\$16.98	\$1.87	-\$0.53	\$5.25	\$2.78	-\$0.87	-\$1.12
2020	\$32.90	\$19.33	\$2.54	-\$0.37	\$5.86	\$3.33	-\$0.75	-\$1.10
2030	\$36.16	\$21.78	\$3.31	-\$0.13	\$6.45	\$3.90	-\$0.55	-\$1.01
2040	\$39.53	\$24.36	\$4.21	\$0.19	\$7.02	\$4.49	-\$0.26	-\$0.82
2050	\$42.98	\$27.06	\$5.25	\$0.63	\$7.53	\$5.09	\$0.14	-\$0.53
% Chg at 2020					-82.2%	-82.8%	-129.5%	-197.3%*

**Table 4:** Mean Social Cost of Carbon estimates by year under four discount rates from the FUND Model, for both the simulated (RB07) and empirical (LC15) ECS distributions. Last row shows the percent change as of 2020. \* Change from -\$0.37 to -\$1.10 is, arithmetically, a positive number, but is shown here as negative to indicate that it is a change to a larger negative magnitude.

Average Standard Deviation – FUND Model								
Discount rates	Using Simulated ECS				Using Empirical ECS			
	2.50%	3.00%	5.00%	7.00%	2.50%	3.00%	5.00%	7.00%
2010	\$64.24	\$31.45	\$5.19	\$2.24	\$67.60	\$42.54	\$7.62	\$2.52
2020	\$32.90	\$35.68	\$6.28	\$2.79	\$80.17	\$52.61	\$10.75	\$3.51
2030	\$36.16	\$40.24	\$7.48	\$3.40	\$93.86	\$64.26	\$15.11	\$5.02
2040	\$39.53	\$45.14	\$8.78	\$4.05	\$108.03	\$77.23	\$21.12	\$7.37
2050	\$42.98	\$50.31	\$10.22	\$4.76	\$121.20	\$90.55	\$29.08	\$11.04
% Chg at 2020					+143.7%	+47.4%	+71.2%	+25.8%

**Table 5:** Average standard deviation of SCC estimates by year under four discount rates from the FUND Model, for both the simulated (RB07) and empirical (LC15) ECS distributions. Last row shows the percent change as of 2020.

Probability of Negative Social Cost of Carbon – FUND Model								
Discount rates	Using Simulated ECS				Using Empirical ECS			
	2.50%	3.00%	5.00%	7.00%	2.50%	3.00%	5.00%	7.00%
2010	0.087	0.121	0.372	0.642	0.416	0.450	0.601	0.730
2020	0.084	0.115	0.344	0.601	0.402	0.432	0.570	0.690
2030	0.080	0.108	0.312	0.555	0.388	0.414	0.536	0.646
2040	0.075	0.101	0.282	0.507	0.371	0.394	0.496	0.597
2050	0.071	0.093	0.251	0.455	0.354	0.372	0.456	0.542

**Table 6:** Probability of a negative Social Cost of Carbon under four discount rates in the FUND Model.